

Combined rotational Raman temperature lidar and water vapor DIAL for humidity, temperature, and particle measurements

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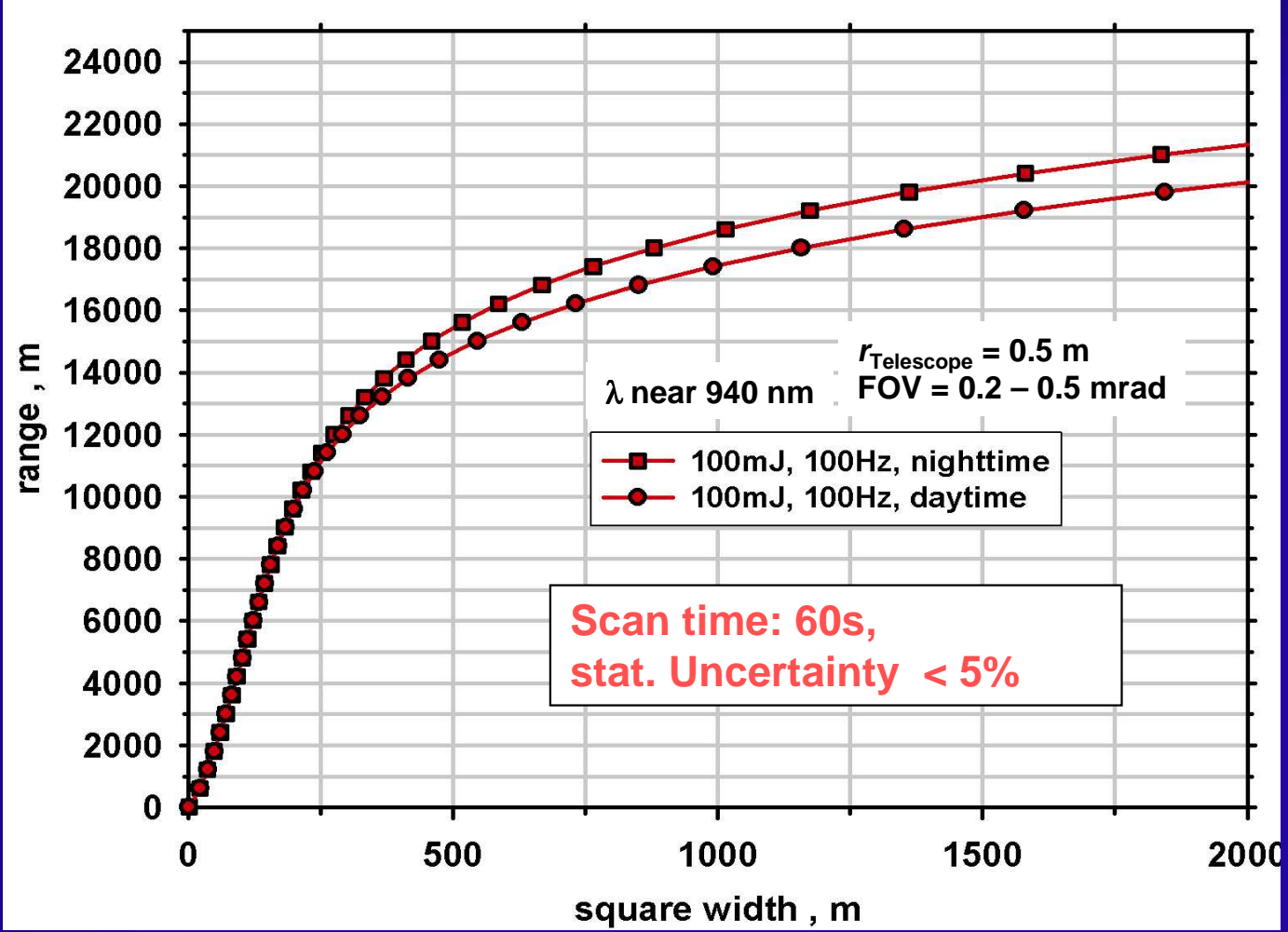
Motivation

**Simultaneous measurement of water vapor and temperature
in planetary boundary layer and free troposphere
with high resolution
and accuracy**

-> RH, CAPE, CIN



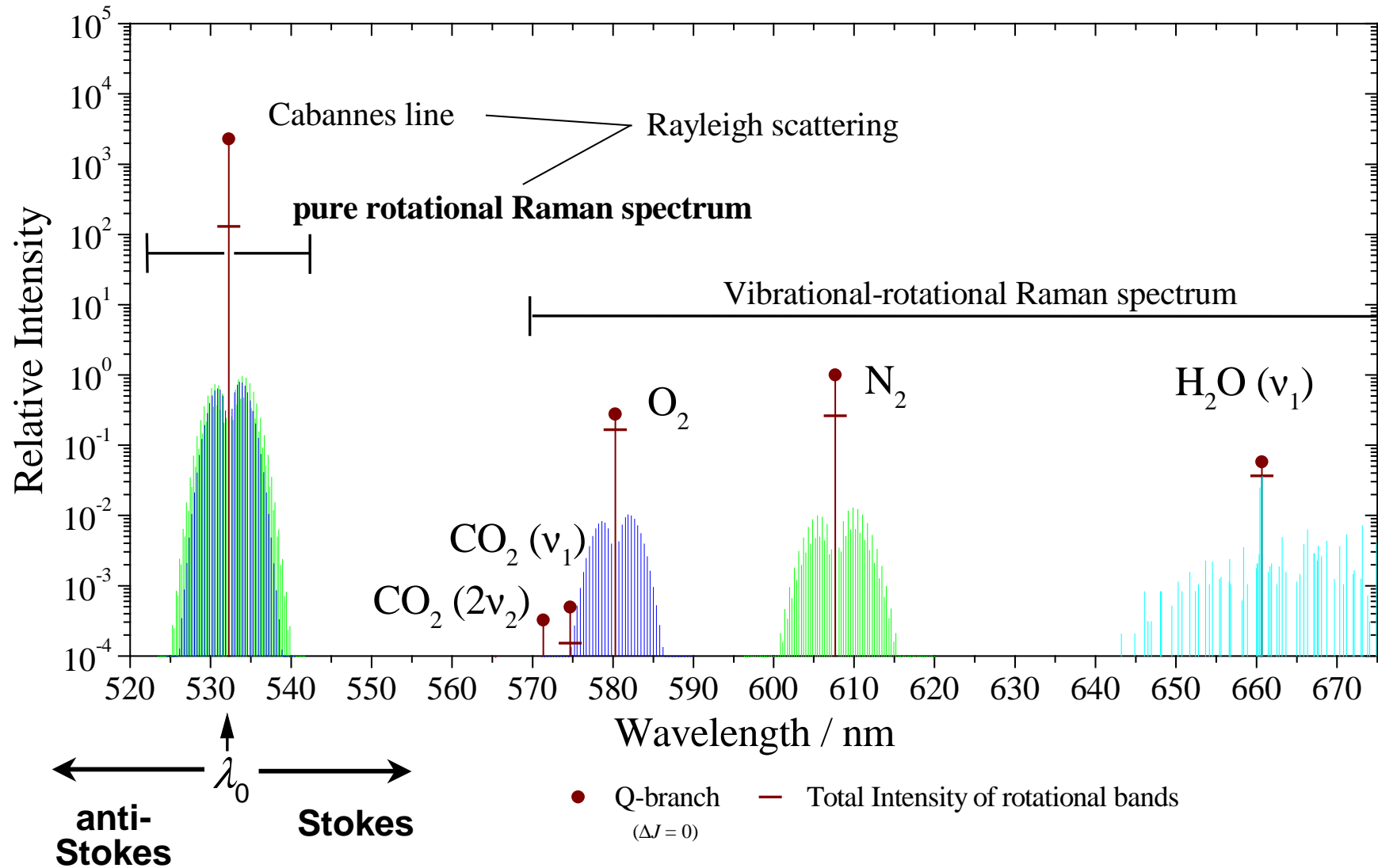
Groundbased scanning H₂O DIAL



Wulfmeyer and
Walther, Appl.
Opt. 2001

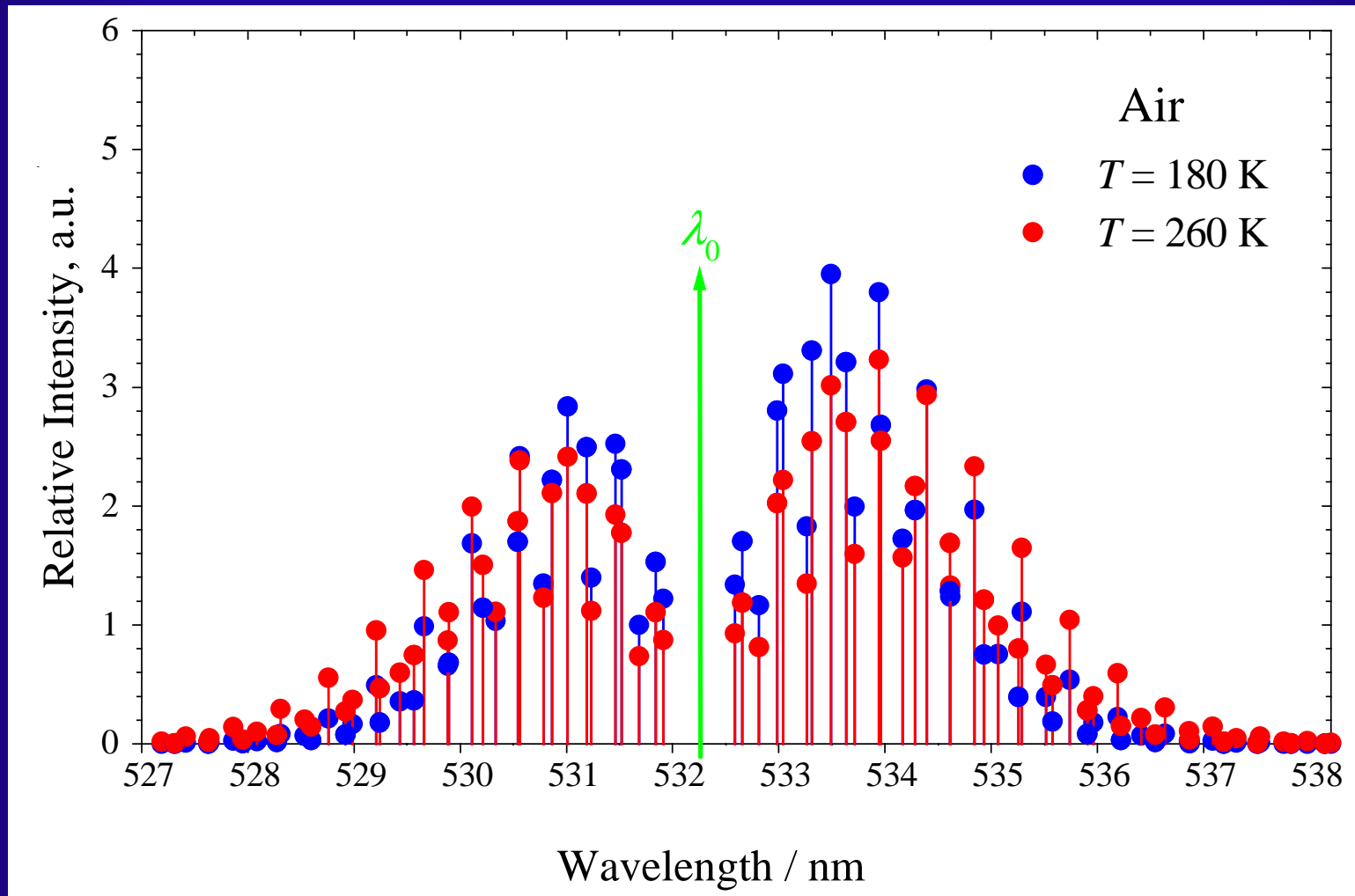


Atmospheric Backscatter Signals

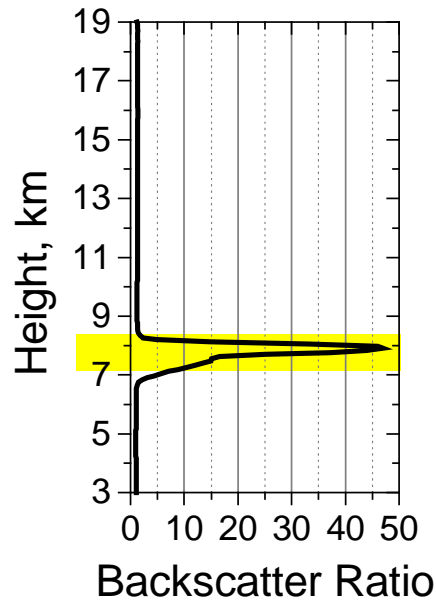
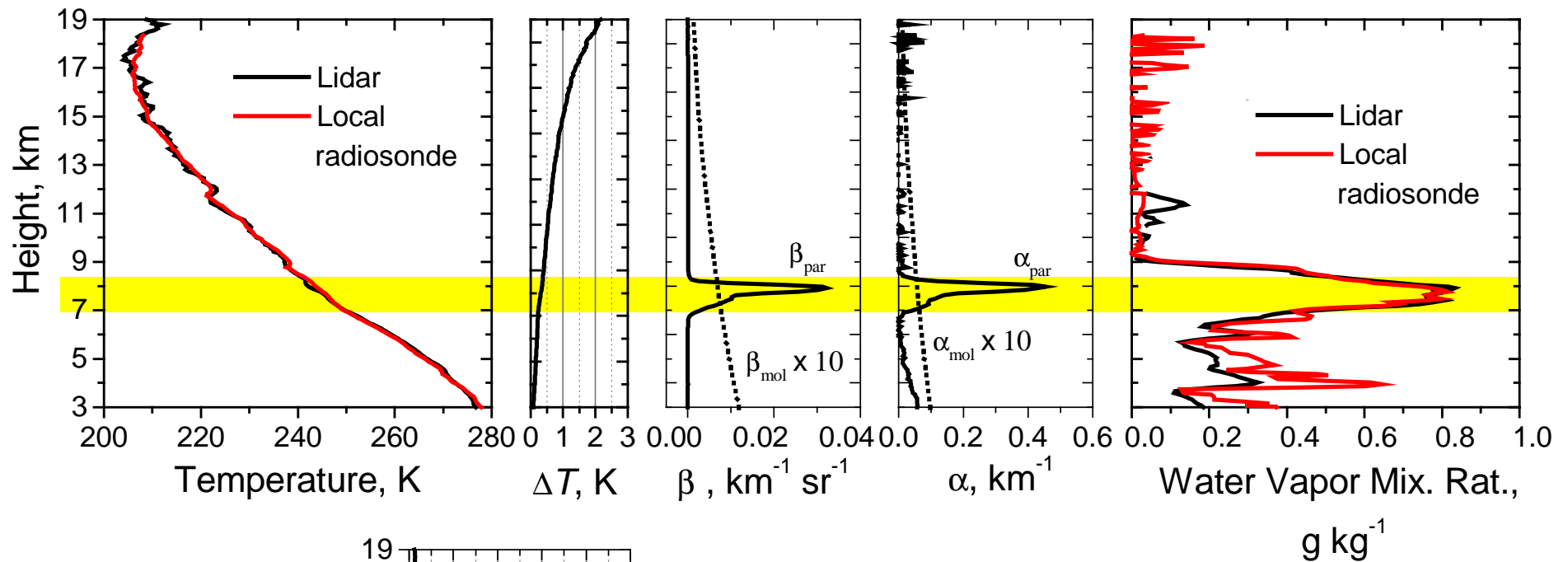


1% H_2O
 $T = 300\text{K}$ (for PRRS)

Rotational Raman Spectrum of Air



$\lambda_0 = 532 \text{ nm}$, 4 signals: Elastic, RR1, RR2, H2O VR



$\lambda = 532 \text{ nm}$
 $P = 30 \text{ W (600 mJ, 50 Hz)}$
 $r_{\text{Telescope}} = 0.5 \text{ m}$
 $\text{FOV} = 1.5 \text{ mrad}$

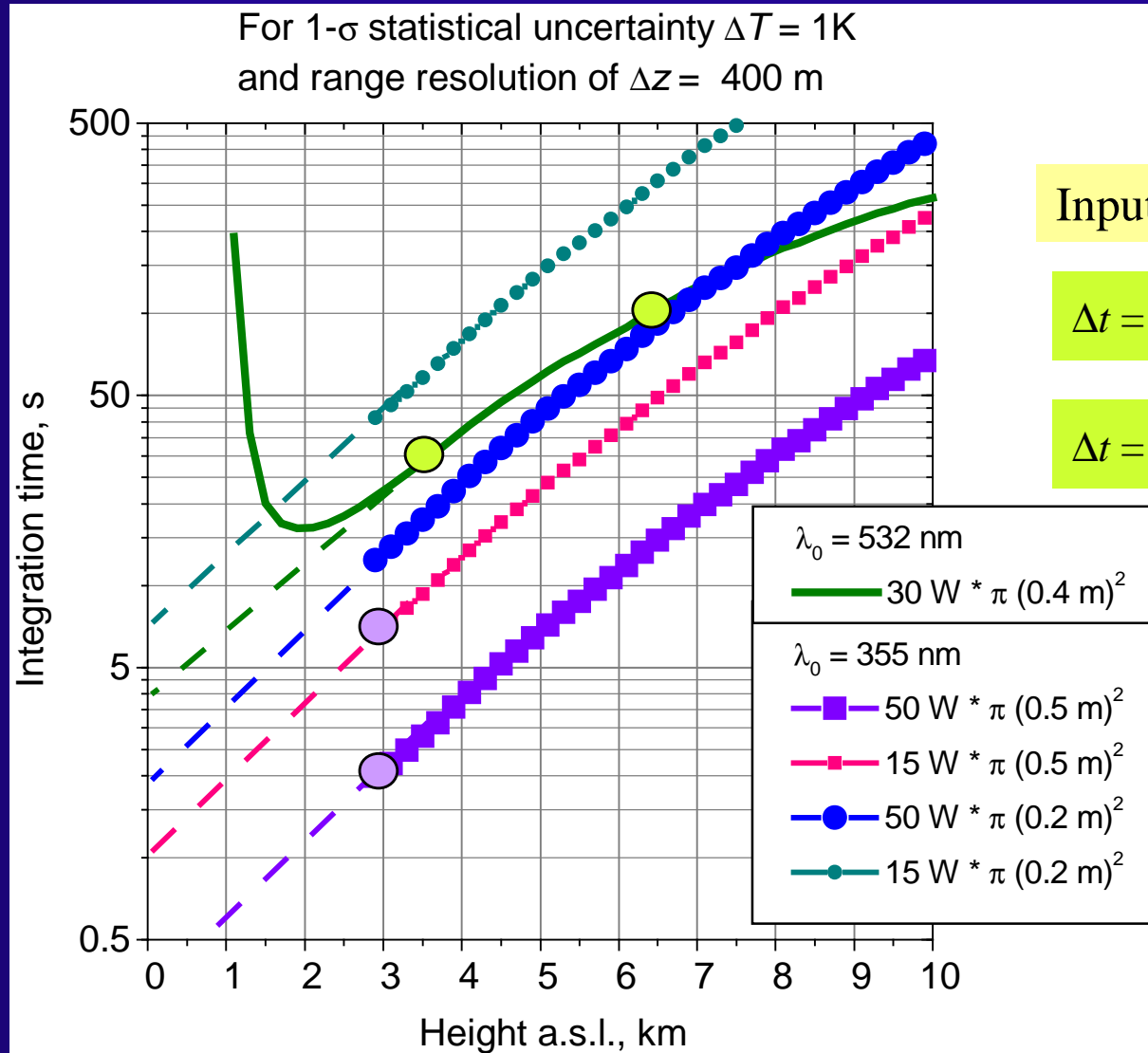
$\Delta t = 30 \text{ minutes,}$
 $\Delta z = 360 \text{ m}$

T measurements with highest resolution
No systematic errors even in thin clouds

A. Behrendt, T. Nakamura, T. Tsuda, Appl. Opt., 2004

Ground-based RR Lidar, vertical profiling

Input: Experimental data of RASC Raman Lidar



Input: RASC Lidar

$\Delta t = 90\text{ s}, z = 6.5\text{ km}$

$\Delta t = 30\text{ s}, z = 3.5\text{ km}$

Future UV Lidar:

$\Delta t = 2 - 7\text{ s}, z = 3\text{ km}$



Theory H₂O DIAL

The water vapor concentration can be found in

$$\ln \left(\frac{\tau_{WV}^2(v_{off}, z_2)}{\tau_{WV}^2(v_{on}, z_2)} \right) - \ln \left(\frac{\tau_{WV}^2(v_{off}, z_1)}{\tau_{WV}^2(v_{on}, z_1)} \right) = 2 \int_{z_1}^{z_2} N_{WV}(r) \sigma_{WV}(r) dr \cong 2 N_{WV,eff} \sigma_{WV,eff} \Delta z$$

but it also shows up in the correction terms!

$$\ln \frac{S_{off}(z_2)}{S_{on}(z_2)} - \ln \frac{S_{off}(z_1)}{S_{on}(z_1)} = \ln \left(\frac{\tau_{WV}^2(v_{off}, z_2)}{\tau_{WV}^2(v_{on}, z_2)} \right) - \ln \left(\frac{\tau_{WV}^2(v_{off}, z_1)}{\tau_{WV}^2(v_{on}, z_1)} \right)$$

$$+ \ln \left[\frac{\beta_{par}(z_2) + \frac{\beta_{mol}(z_2) \int_{v=0}^{\infty} b(v - v_{off}) \tau_{WV}(v, z_2) dv}{\tau_{WV}(v_{off}, z_2)}}{\beta_{par}(z_2) + \frac{\beta_{mol}(z_2) \int_{v=0}^{\infty} b(v - v_{on}) \tau_{WV}(v, z_2) dv}{\tau_{WV}(v_{on}, z_2)}} \right] - \ln \left[\frac{\beta_{par}(z_1) + \frac{\beta_{mol}(z_1) \int_{v=0}^{\infty} b(v - v_{off}) \tau_{WV}(v, z_1) dv}{\tau_{WV}(v_{off}, z_1)}}{\beta_{par}(z_1) + \frac{\beta_{mol}(z_1) \int_{v=0}^{\infty} b(v - v_{on}) \tau_{WV}(v, z_1) dv}{\tau_{WV}(v_{on}, z_1)}} \right]$$

H₂O DIAL

“Simulation results show that great care has to be taken in the analysis of H₂O DIAL measurements when **layers with high aerosol concentration, clouds, or strong temperature inversion** exist.”

(Ansmann, AO 1985)

To correct for errors caused the Rayleigh-Doppler effect, information on the aerosol backscattering properties is necessary.

These can be retrieved from the off-line signal.

For this retrieval **Rayleigh scattering coefficients, aerosol extinction/backscatter ratio, and the aerosol backscatter coefficient** at a calibration range have to be known. (Klett-algorithm)

(Ansmann & Boesenberg, AO 1987)



Concept

Idea: Correction of Rayleigh-Doppler effect by use of rotational Raman signal measured in the UV.

Scheme:

- 1) Measurement of $\alpha_{\text{par,IR}}, \beta_{\text{par,IR}}$ with lower resolution
 $\alpha_{\text{par,UV}}, \beta_{\text{par,UV}}$ with high resolution



- 2) Radiative properties of atmospheric particles constant in time scale of minutes (concentration of particles not necessarily constant)

$$\frac{\alpha_{\text{par,IR}}}{\alpha_{\text{par,UV}}} = \text{constant} \quad \frac{\beta_{\text{par,IR}}}{\beta_{\text{par,UV}}} = \text{constant} \quad \alpha_{\text{par,UV}}, \beta_{\text{par,UV}} \Rightarrow \alpha'_{\text{par,IR}} \text{ and } \beta'_{\text{par,IR}} \text{ with high resolution}$$

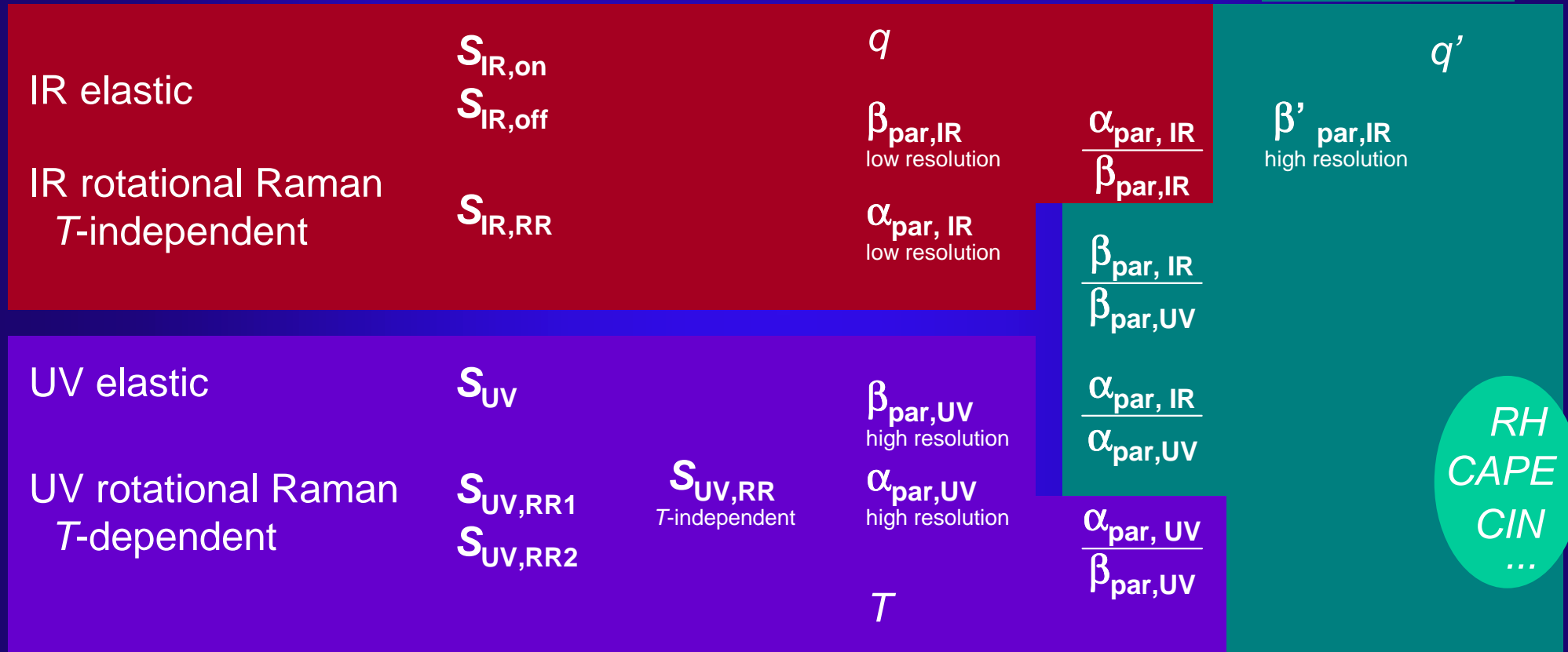


- 3) Use measured $\alpha'_{\text{par,IR}}, \beta'_{\text{par,IR}}, T$ to correct for Rayleigh-Doppler effect



Possible scheme

Synergy



RH
CAPE
CIN
...

Behrendt and Wulfmeyer, SPIE 2003.

